

中国东部中生代钼矿成岩-成矿时差

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摘要:与岩浆成因相关的钼矿床成岩-成矿时差是确定矿床与岩浆岩成因关系的重要基础。本文根据收集的中国东部中生代典型钼矿床的成矿及相关岩体的同位素测年数据, 详细讨论并定量厘定了钼矿的成岩-成矿时差分布特征。结果表明, 钼矿成矿同步或略滞后于同源岩浆活动, 中国东部整个钼矿成矿高峰的两个阶段时差介于 0~10.0 Ma 和 0~15.0 Ma; 对于单个钼矿床, 其成岩-成矿时差集中在 0~14.0 Ma, 均值为 3.9 Ma; 从斑岩型钼矿床→斑岩-矽卡岩型钼矿床→矽卡岩型钼矿床→石英脉型钼矿床, 成岩-成矿时差呈逐渐增加趋势, 这与岩浆热液成矿过程的地质事实吻合。

关键词:钼矿; 成岩-成矿时差; 中国东部; 中生代

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A Study of the Time Gap between Diagenesis-Mineralization of Mesozoic Molybdenum Deposits in East China

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Abstract: The time gap of diagenesis-mineralization, a hotspot in Mo deposits study, is an important basis to confirm the genetic relationship between Mo deposits and their cognatemagmatites. According to isotopic dating of mineralization and associated rock of some typical Mesozoic Mo deposits in East China, we detailly discusse and quantitatively determinate the time gap of diagenesis-mineralization for Mo deposits. Mo mineralization occurred synchronously or slightly behind the comagmatic activity. Two peak stages of diagenesis- mineralization time gap of Mo mineralization are within the scope of 0-10.0 Ma and 0-15.0 Ma; For a single Mo deposit, the time gap concentrated in 0-14.0Ma with a mean of 3.9 Ma. From the porphyry-type Mo deposits to porphyry-skarn complex Mo deposits to skarn-type Mo deposits and to quartz vein-type Mo deposits, the time gap of diagenesis and mineralization increases gradually which is consistent with the objective geological facts of the magmatic hydrothermal ore-forming process.

Key words: Mo deposits; time gap of diagenesis and mineralization; East China; Mesozoic

中国的钼矿储量位居世界第二, 主要集中在东部, 仅东秦岭的钼矿储量就约占全国总储量的 43.5%。钼矿床有斑岩型、斑岩-矽卡岩复合型、矽卡岩型、石英脉型、爆破角砾岩型、碳酸岩脉型、沉积型和伟晶岩脉型等类型, 以斑岩型、斑岩-矽卡岩复合型最重要, 其次为矽卡岩型和石英脉型, 这四种类型钼矿床成矿都与岩浆活动密切相关^[1,2]。从钼矿形成时代来看, 除少数钼矿床形成于元古宙、晚古生

代和新生代之外, 绝大多数都形成于中生代^[2~4]。

钼矿成岩-成矿时差, 是研究矿床成因、指导找矿以及建立区域成矿模式的重要基础, 然而目前这方面的研究还比较少。研究整个钼矿成矿阶段及单个矿床和与其有成因关系的岩浆岩的成岩-成矿时差, 对确定岩浆演化与成矿的关系不仅具有理论意义, 更具有指导找矿的实践意义。

近几年我国东部地区钼矿研究出现了新一轮高

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潮,找矿勘查有了新发现和新进展,高精度同位素技术精确厘定成岩-成矿时代,为从时差角度研究钼矿成因提供了重要素材。本文根据收集的近十几年来中国东部典型钼矿床的成矿年龄及同源岩体成岩年龄数据,讨论了钼矿床的成岩-成矿时差特征。

1 钼矿成矿背景

中国东部大地构造演化大致经历了太古宙—古元古代原始地壳克拉通化阶段,中、新元古代—早古生代大陆边缘或陆内盆地演化阶段,晚古生代欧亚板块拼合,古欧亚大陆形成阶段和中、新生代受太平洋板块和印度板块俯冲产生的叠加构造阶段,构成了滨太平洋成矿域^[5,6,7]。从三叠纪开始,古太平洋板块与初始欧亚板块逐渐发生挤压碰撞与俯冲作用^[8],燕山期时中国东部处于大陆边缘活动带,俯冲作用引起了地壳内部及壳—幔物质交换,同时形成

了交代型富集地幔^[9],演化晚期产生了与壳幔同熔型-地壳重熔型酸性-中酸性岩浆岩有关的中生代钼矿成矿带,如东秦岭地区、燕辽地区、长江中下游地区和南岭地区等(图 1)。从钼矿形成时代来看,主要在 190.0~110.0 Ma(图 2),为中生代燕山期构造岩浆活动的产物。

2 钼矿成岩-成矿时差

2.1 数据收集

本文钼矿床成岩-成矿年龄数据主要为与岩浆活动密切相关的斑岩型、斑岩-矽卡岩型、矽卡岩型和石英脉型等钼矿床。由于 Re-Os 同位素体系封闭性好,受后期改造微弱^[10],被广泛应用于同位素地质年代学、矿床成因等研究中,尤其对钼矿床测年,国内外将辉钼矿 Re-Os 法作为最主要的成矿测试方法。文中数据除了个别矿床采用石英流体包裹体⁴⁰Ar/³⁹Ar 法,绝大部分为该方法测得(40 个/43 个)。对与成矿有关的花岗质岩石年龄,采用 SHRIMP 锆石 U-Pb 法、单颗粒锆石 U-Pb 法(27 个/43 个)和⁴⁰Ar/³⁹Ar 法及其它方法(见表 1)。

2.2 数据筛选

与钼矿床成矿有关的同源岩体有两种情况:一是通过岩浆期后分异作用提供成矿流体和成矿物质的花岗质岩类;二是与成矿流体和成矿物质均来自深部的同一岩浆区的花岗质岩类^[73]。

首先选取典型矿床,并对其成岩-成矿的关系进行判断是否同源。如:卢欣祥等^[74]、黄典豪等^[75]、朱赖民等^[37]通过对东秦岭金堆城超大型钼矿床的岩矿石微量元素和稀土元素分布特征以及 Pb、S、Sr、O 等同位素组成对比研究,结合成岩-成矿年代,认为钼成矿与花岗斑岩侵入直接相关;研究吉林大黑山超大型钼矿床岩矿石稀土配分曲线、S 同位素、Re 含量和成岩-成矿年龄发现,钼成矿与花岗闪长斑岩密切相关^[14~15,76];孟良义等^[77]、黄恩邦等^[78]、罗建安等^[79]相继根据岩体与矿体的接触关系和空间产状、成矿元素的分布特征、黄铁矿的 Co/Ni 和 S/Se、Sr、Pb 及 S、H、O、C 等稳定同位素对比,均认为城门山钼矿床的成矿物质和成矿流体均来源于岩浆岩,成岩与成矿物质具有同源性。

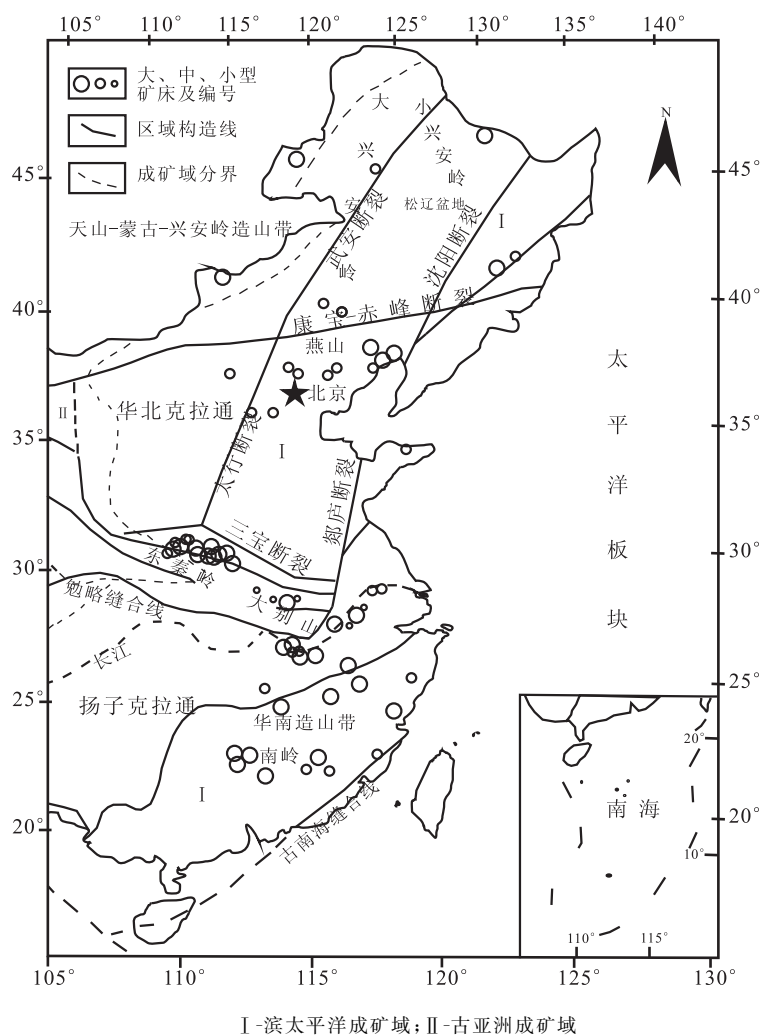


图 1 中国东部钼矿床分布图

Fig. 1 The distribution map of Mo deposits in the eastern China

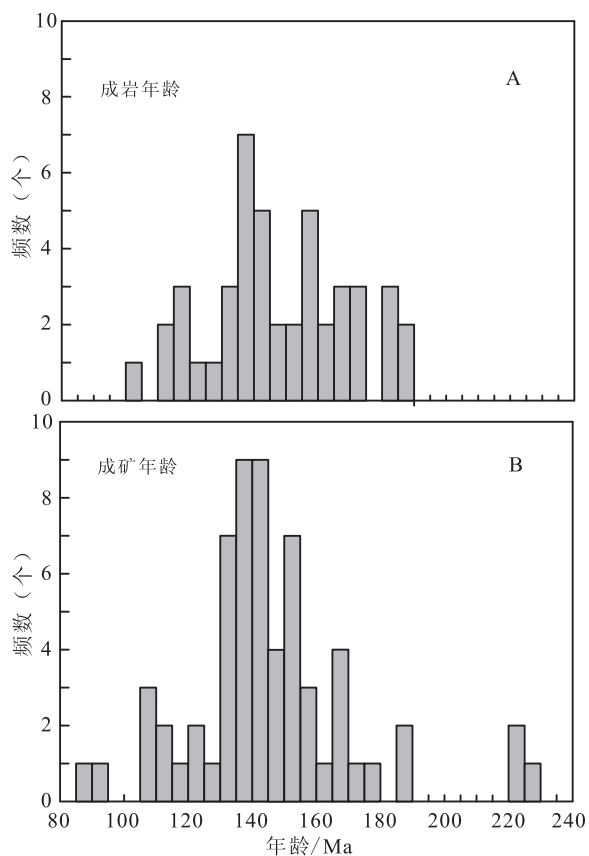


图2 中国东部中生代钼矿床成岩年龄(A)、成矿年龄(B)分布直方图

Fig. 2 Histogram of isotopic ages of diagenesis (A) and mineralization (B) for Mesozoic Mo deposits in East China

其次,数据收集时也考虑了成矿时间的可比性、不同测年方法的精确性和区域成矿背景,选择研究较深入且被公认可靠性较高的年龄值。如辽西兰家沟大型钼矿床全岩 Rb-Sr 法测得成岩年龄 154 Ma^[79],而锆石 SHRIMP U-Pb 法测得成岩年龄 188.9 Ma^[22],辉钼矿 Re-Os 法测得成矿年龄 186.5 Ma^[3],由于成岩年龄更接近成矿年龄,因此认为其成岩年龄为 188.9 Ma。同样豫西雷门沟超大型钼矿床, K-Ar 法获得成岩年龄为 99.3~88.4 Ma^[81]; 锆石 SHRIMP U-Pb 法测得成岩年龄为 136.2 Ma, 辉钼矿 Re-Os 法测得成矿年龄为 132.4 Ma^[32], 东秦岭钼矿大规模成矿于晚侏罗世和早白垩世之交的地质背景一致,采用 136.2 Ma 更加符合事实。

另外考虑到个别钼矿床时差太大,不符合地质事实,如内蒙古乌奴格吐山斑岩铜钼矿床(28 Ma),河北寿王坟铜(钼)矿床(-18 Ma),应预先剔除。

2.3 数据统计

图2可知,中国东部中生代与钼成矿相关的成岩年龄集中在 160~155 Ma 和 145~135 Ma, 相应成矿年龄集中在 155~150 Ma 和 145~130 Ma,反

映出中国东部钼矿成矿具有两次高峰期。第一阶段成岩-成矿的总体时差为 0~10 Ma, 第二阶段成岩-成矿的总体时差为 0~15 Ma。

单个钼矿床的成岩-成矿时差数据表明(表1),钼矿床的成矿时间同步或略滞后于同源岩浆活动,平均时差介于零到十几个百万年。另外,部分钼矿床的成岩-成矿时差为负值(即成矿早于成岩),如:东秦岭东沟钼矿床(-4 Ma)、内蒙古乌兰德勒钼铜多金属矿床(-2.7 Ma),不具实际地质意义,可能由测量方法、测试误差或测试对象的同位素封闭程度导致。由图3可见,钼矿床成岩-成矿时差一般不超过 16 Ma,集中在 0~14 Ma。由图4也可见单个钼矿床与同源岩体的时差集中于 0~14 Ma,时差平均值为 3.9 Ma,具有正态分布变量的成因特征,符合统计意义。

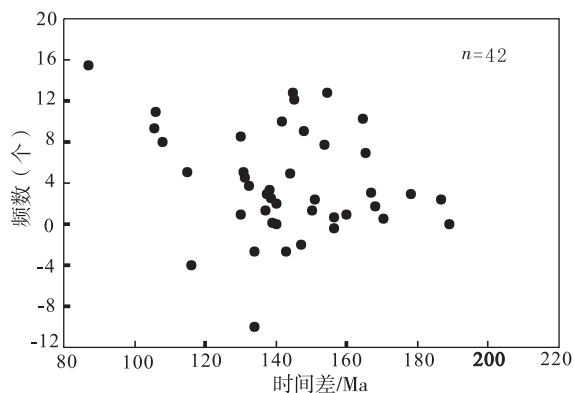


图3 钼矿床成岩成矿时差(数据来自表1)

Fig. 3 Diagram of diagenesis-mineralization time gap and mineralization ages for comagmatic Mo deposits (Data from Table 1)

笔者还针对不同类型的钼矿床成岩-成矿时差进行了初步统计(表1),时差平均值明显不同。斑岩型钼矿床时差平均值为 2.6 Ma;斑岩-矽卡岩型(或矽卡岩-斑岩型)钼矿床时差平均值为 3.8 Ma;矽卡岩型钼矿床时差平均值为 6.6 Ma;石英脉型钼矿床时差平均值为 8.6 Ma。

2.4 时差讨论

根据以上统计,中国东部钼矿成矿高峰的两个阶段成岩-成矿时差介于 0~10.0 Ma 和 0~15.0 Ma。对于单个钼矿床,则将与岩浆密切相关的钼矿床成岩-成矿时差限定在 0~14.0 Ma,平均约为 3.9 Ma。尽管也有学者认为,单一侵入体所引起的热液活动最长时间小于 1.0 Ma^[82]。在目前的定年测试条件下,笔者认为,成岩-成矿年龄要精确到 1.0 Ma 还是很困难的,而且在地质问题的讨论上也没有太大的实际意义。

表 1 中国东部中生代钼矿床及其同源岩浆岩同位素年龄
 Table 1 Isotope ages of Mesozoic Mo deposits and related comagmatic magmatites in East China

编号	矿床名称	矿床类型	同源岩浆岩				矿床				时差 /Ma
			岩性	测定对象	测定方法	年龄/Ma	测定对象	测定方法	年龄/Ma	资料来源	
1	内蒙古太平沟钼矿床	斑岩型	花岗斑岩	锆石	SHRIMP U-Pb	131.1±0.9	辉钼矿	Re-Os	130.1±1.3	文献[11]	1
2	内蒙古乌奴格吐山铜钼矿床	斑岩型	二长花岗斑岩	锆石	U-Pb	183±0.6	辉钼矿	Re-Os	155±17	文献[12],[13]	28
3	吉林大黑山钼矿床	斑岩型	花岗闪长斑岩	锆石	SHRIMP U-Pb	170±3	辉钼矿	Re-Os	168.2±3.2	文献[14],[15]	1.8
4	吉林福安堡钼矿床	斑岩型	似斑状二长花岗岩	全岩	Rb-Sr	170	辉钼矿	Re-Os	166.9±6.7	文献[16],[17]	3.1
5	内蒙古碾子沟钼矿床	石英脉型	黑云母二长花岗岩	全岩	Rb-Sr	167.1±1.5	辉钼矿	Re-Os	154.3±3.6	文献[18],[19]	12.8
6	内蒙古乌兰德勒钼矿床	斑岩型	细粒二长花岗岩	锆石	SHRIMP U-Pb	131.3±1.6	辉钼矿	Re-Os	134.0±4.2	文献[20]	-2.7
7	辽西杨家杖子钼矿床	斑岩-矽卡岩型	细粒花岗岩	锆石	U-Pb	189±4	辉钼矿	Re-Os	189	文献[21],[3]	0
8	辽宁兰家沟钼矿床	斑岩型	似斑状细粒花岗岩	锆石	SHRIMP U-Pb	188.9±1.2	辉钼矿	Re-Os	186.5±0.7	文献[22],[3]	2.4
9	辽西新台门钼矿床	斑岩型	花岗岩	锆石	LA-ICP-MS U-Pb	181±2	辉钼矿	Re-Os	178±5	文献[23]	3
10	河北小寺沟钼(铜)矿床	斑岩-矽卡岩型	花岗闪长斑岩	全岩	Rb-Sr	124±1.9	辉钼矿	Re-Os	134±3	文献[24],[3]	-10
11	河北大草坪钼矿床	斑岩型	花岗闪长岩	锆石	U-Pb	140.0±1.5	辉钼矿	Re-Os	140.0±2.3	文献[25]	0
12	河北寿王坟钼(铜)矿床	矽卡岩型	花岗闪长岩	全岩	Rb-Sr	130	辉钼矿	Re-Os	148±4	文献[26],[3]	-18
13	河南南泥湖钼矿床	斑岩型	斑状二长花岗岩	锆石	SHRIMP U-Pb	157.1±2.9	辉钼矿	Re-Os	148±10	文献[27],[28]	9.1
14	河南秋树湾钼钨矿床	斑岩-矽卡岩型	花岗闪长斑岩	黑云母	K-Ar	145	辉钼矿	Re-Os	147±4	文献[29],[30]	-2
15	河南三道庄钼矿床	矽卡岩型	斑状二长花岗岩	锆石	SHRIMP U-Pb	157.1±2.9	辉钼矿	Re-Os	145.0±2.2	文献[27],[31]	12.1
16	河南上房沟钼矿床	斑岩-矽卡岩型	花岗斑岩	锆石	SHRIMP U-Pb	157.6±2.7	辉钼矿	Re-Os	144.8±2.1	文献[27],[31]	12.8
17	河南雷门沟钼矿床	斑岩型	花岗斑岩	锆石	SHRIMP U-Pb	136.2±1.5	辉钼矿	Re-Os	132.4±2.0	文献[32]	3.8
18	河南鱼池岭钼矿床	斑岩型	似斑状二长花岗岩	黑云母	Ar-Ar	135.7±1.3(1)	辉钼矿	Re-Os	131.2±1.4	文献[33],[34]	4.5
19	河南泉家峪钼金矿床	石英脉型	花岗岩	锆石	SHRIMP U-Pb	138.4±2.5	辉钼矿	Re-Os	129.9±1.6	文献[27],[35]	8.5
20	河南东沟钼矿床	斑岩型	花岗斑岩	锆石	SHRIMP U-Pb	112±1	辉钼矿	Re-Os	116±1.7	文献[36]	-4
21	陕西金堆城钼矿床	斑岩型	花岗斑岩	锆石	LA-ICP-MS U-Pb	141.0±0.5	辉钼矿	Re-Os	138.4±0.5	文献[37],[38]	2.6
22	陕西八里坡钼矿床	斑岩型	花岗斑岩	锆石	U-Pb	155.9±2.3	辉钼矿	Re-Os	156.3±2.2	文献[39]	-0.4

续表 1

编号	矿床名称	矿床类型	同源岩浆岩			矿床			时差 /Ma		
			岩性	测定对象	测定方法	年龄/Ma	测定对象	测定方法		年龄/Ma	
23	陕西石家湾钨矿床	斑岩型	钾长花岗岩斑岩	锆石	LA-ICP-MS U-Pb	141.4±0.6	辉钨矿	Re-Os	138	文献[40]、[28]	3.4
24	河南汤家坪钨矿床	斑岩型	花岗岩	锆石	U-Pb	120±3	辉钨矿	Re-Os	114.9±2.7	文献[41]、[42]	5.1
25	湖北丰山洞铜钨矿床	矽卡岩-斑岩型	花岗闪长斑岩	黑云母	K-Ar	149	辉钨矿	Re-Os	144.0±2.1	文献[43]、[44]	5
26	湖北铜绿山铜钨矿床	矽卡岩型	花岗闪长岩	锆石	SHRIMP U-Pb	140.3±1.8	辉钨矿	Re-Os	137.3±2.4	文献[45]	3
27	湖北铜山口铜钨矿床	矽卡岩-斑岩型	花岗闪长斑岩	锆石	SHRIMP U-Pb	140.2±2.4	辉钨矿	Re-Os	142.9±1.8	文献[46]、[44]	-2.7
28	安徽铜牛井铜钨矿床	石英脉型	闪长岩	角闪石	Ar-Ar	136±1.8 (p)	石英	Ar-Ar	130.9±0.8 (d)	文献[47]、[48]	5.1
29	安徽铜陵沙滩角铜钨矿床	矽卡岩型	石英二长斑岩	锆石	SHRIMP U-Pb	151.8±2.6	辉钨矿	Re-Os	141.8±1.6	文献[49]、[27]	10
30	安徽大团山铜钨矿床	矽卡岩型	石英二长闪长岩	锆石	SHRIMP U-Pb	139.3±1.2	辉钨矿	Re-Os	139.1±2.7	文献[50]、[51]	0.2
31	江苏安基山铜钨矿床	矽卡岩-斑岩型	花岗闪长斑岩	黑云母	K-Ar	116	辉钨矿	Re-Os	108±2	文献[52]	8
32	江苏铜山铜钨矿床	矽卡岩型	石英闪长玢岩	黑云母	K-Ar	117	辉钨矿	Re-Os	106±3	文献[52]	11
33	江西德兴铜钨矿床	斑岩型	花岗闪长斑岩	锆石	SHRIMP U-Pb	171±3	辉钨矿	Re-Os	170.4±1.8	文献[53]、[54]	0.6
34	江西城门山铜钨矿床	矽卡岩-斑岩型	石英斑岩	黑云母	K-Ar	142	辉钨矿	Re-Os	140±2	文献[55]	2
35	江西阳储岭钨矿床	斑岩型	二长花岗岩	全岩	Rb-Sr	138.4±3.4	白云母	K-Ar	137.1±3.4	文献[56]	1.3
36	江西浒坑钨钼矿床	石英脉型	白云母花岗岩	锆石	LA-ICP-MS U-Pb	151.6±2.6	辉钨矿	Re-Os	150.2±2.2	文献[57]、[58]	1.4
37	湖南宝山铜钨多金属矿床	矽卡岩-斑岩型	花岗闪长斑岩	锆石	SHRIMP U-Pb	161±3	辉钨矿	Re-Os	160±2	文献[59]、[60]	1
38	湖南黄沙坪铅锌钨矿床	矽卡岩型	花岗斑岩	锆石	LA-ICP-MS U-Pb	161.6±1.1	辉钨矿	Re-Os	153.8±4.8	文献[61]、[62]	7.8
39	湖南柿竹园钨钼钨矿床	矽卡岩型	中粗粒钾长花岗岩	黑云母	Ar-Ar	153.4±0.2	辉钨矿	Re-Os	151±3.5	文献[63]、[64]	2.4
40	浙江石平川钨矿床	石英脉型	钾长花岗岩	锆石	LA-ICP-MS U-Pb	102.5±1.2	石英	Rb-Sr	87±1	文献[65]、[66]	15.5
41	福建行洛坑钨钼矿床	斑岩型	花岗岩	全岩	Rb-Sr	157±3	辉钨矿	Re-Os	156.3±4.8	文献[67]、[68]	0.7
42	福建安南路钨钼矿床	斑岩型	似斑状花岗岩	全岩	Rb-Sr	115±4	辉钨矿	Re-Os	105.6±0.6	文献[69]、[70]	9.4
43	福建永定山口钨矿床	斑岩型	中细粒二长花岗岩	全岩	Rb-Sr	172.3±5.6	辉钨矿	Re-Os	165.3±3.5	文献[71]	7
44	广东大宝山铜钨矿床	斑岩-矽卡岩型	花岗闪长斑岩	锆石	LA-ICP-MS U-Pb	175.0±1.7	辉钨矿	Re-Os	164.7±3	文献[72]、[63]	10.3

注:表中部分数据按有效数字规范进行了取舍(保留一位小数),部分年龄未标出误差;对于同一成岩或成矿年龄的不同测试方法采用更新更精确的方法;p,1分别为 Ar-Ar法中坪年龄和等时线年龄。

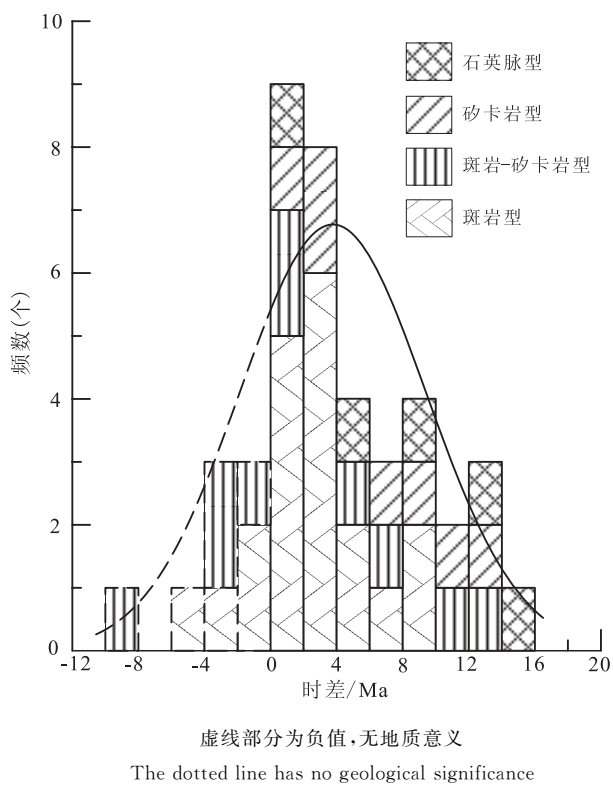


图4 钼矿床成岩成矿时差直方图(数据来自表1)

Fig. 4 Histogram of diagenesis-mineralization time gap for comagmatic molybdenum deposits (Data from Table 1)

另外,考虑到钼矿化一般与花岗岩类小岩株有关,对于一个 30 km^2 的隐伏岩体(埋深在 $1 \sim 2 \text{ km}$ 以下),从 800°C 的岩浆冷却到岩体中心为 600°C 时,冷凝时间通常仅需 $7.2 \sim 10.0 \text{ Ma}$ ^[83, 84];同时一个大型、超大型钼矿床的形成,往往具有多期次成矿特征,即在上部小岩株侵位、冷凝、收缩产生裂隙的同时,由于构造-脉动作用使之与下部岩浆联系一直未中断,岩浆房继续分异,同时岩体与围岩接触带构造断裂加剧,从而成为含钼热流体源源不断上升的通道,促成大型、超大型钼矿床各阶段各类矿化的发展、叠加、富集。如东秦岭的南泥湖和上房沟两个超大型钼矿床的成岩成矿时差较大($9 \sim 14 \text{ Ma}$),笔者认为有两种可能,一是一个超大型钼矿床的形成需要多期次多阶段的岩浆活动,持续时间较长;二是从测试的辉钼矿采样位置看,南泥湖为条带状角岩,上房沟为接触带外侧^[32],都应是岩浆热液运移后期作用形成的辉钼矿,测年反映的时间会偏晚些。

钼矿化温度($240 \sim 450^\circ\text{C}$)低于花岗岩类冷凝封闭温度(一般大于 600°C),说明钼矿床中同源岩体与矿床形成的年龄差是客观存在的。它的形成受控于岩浆侵位的深度、岩体大小、岩浆成分、地热梯度、构造、围岩性质等,是许多随机因素综合作用的结

果,因此时差获取的关键还是成岩-成矿年龄的准确测定及对矿床成因的正确把握。文中收集的成矿年龄绝大多数是通过辉钼矿 Re-Os 法测得,成岩年龄多采用锆石 U-Pb 法以及 $^{40}\text{Ar}/^{39}\text{Ar}$ 法,运用这些方法得到的就是直接的成岩-成矿年龄,二者差值反映更接近于矿床的真时差。另外笔者还综合考虑了岩体与成矿的同源性、区域地质背景和成矿时间的可比性及单个矿床年代学数据的可靠性,从多方面来厘定同源岩浆成因钼矿床的时差。

大型超大型同源岩浆钼矿床的成岩-成矿时差可能达到 14.0 Ma 。在同一构造背景条件下,若成岩与成矿存在着小于等于 14.0 Ma 时差,就应该重视二者之间的联系(岩浆期后热液成矿与深部同源成矿均有可能),并结合微量元素和同位素示踪进一步揭示其成因关系;而当成岩与成矿存在着大于 14.0 Ma 时差时,就可以排除矿床与岩体的直接成因联系,只能将岩体作为容矿围岩,但其也可能提供部分成矿物质。

谭俊等^[73]曾测得金矿床成岩-成矿时差平均值,为 7.4 Ma ,而此次获得的钼矿床成岩-成矿时差平均值为 3.9 Ma ,很可能是钼矿化温度一般属于高温和中高温,且在岩体内部或接触带,而金矿化温度属于中低温,且多在岩体外侧或围岩裂隙中,成矿热液冷却沉淀成矿的时间不同所造成的,当然不排除现今采用了更为先进的测年手段和文中选用的精确数据较多的因素影响。

本文对不同类型钼矿床成岩-成矿时差进行的初步统计显示,斑岩型、斑岩-矽卡岩型、矽卡岩型、石英脉型钼矿床时差平均值依次为 2.6 Ma 、 3.8 Ma 、 6.6 Ma 、 8.6 Ma ,表明从斑岩型钼矿床→斑岩-矽卡岩复合型钼矿床→矽卡岩型钼矿床→石英脉型钼矿床,成岩-成矿时差呈逐渐增加趋势,这恰与岩浆热液从岩体内部→岩体与围岩接触带→围岩裂隙的迁移富集成矿过程的客观地质事实相吻合。因此从不同类型钼矿床总结出的时差规律有望作为钼矿床类型的一个重要参考指标,并对于确定钼矿床成因及指导进一步找矿勘查亦具有重要的指示意义。然而至今各类型钼矿床的同位素数据还不够充分,这个规律还有待进一步验证。当然,随着同位素地质年代学的进一步完善,钼矿床研究的深入及更多高精度数据的出现,有关钼矿床成岩-成矿时差的厘定将会更加具体,从而确定出更为符合客观事实的时差规律,这将对相关矿床的成因探讨,诸如年代学和地球化学数据的解释等均有指导作用。

3 结 论

(1) 钼成矿同步或略滞后于同源岩浆活动, 中国东部整个钼矿成矿高峰的两个阶段成岩-成矿时差为 0~10.0 Ma 和 0~15.0 Ma; 单个钼矿床成岩-成矿时差集中在 0~14.0 Ma, 平均 3.9 Ma;

(2) 从斑岩型钼矿床→斑岩-矽卡岩型钼矿床→矽卡岩型钼矿床→石英脉型钼矿床, 成岩-成矿时差呈逐渐增加趋势, 这恰与岩浆热液成矿过程的客观地质事实相吻合。

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